

## Introduction

Fire frequency is increasing rapidly in the Arctic, and fires are a significant source of greenhouse gas emissions. Better accounting of emissions through combustion from fires is necessary for understanding their effect on the climate. Estimating the effect of greenhouse gases from remote fires is difficult because it requires measuring the biomass and carbon loss *in situ* in order to calculate radiative forcing.

We studied a 2015 burn scar in the Yukon-Kuskokwim Delta, AK to identify potential carbon loss and radiative forcing from a tundra fire using a new method to estimate organic matter loss with plant markers and temporally-explicit radiative forcings of gaseous emissions.

## Methodology **Estimate biomass and carbon loss**

Two genera of moss, *Dicranum* and *Sphagnum*, and tussock grass, and N<sub>2</sub>O *Eriophorum vaginatum*, survive the fire and indicate pre fire soil level

Equation 1. Burn depth equation based on heights of vegetation markers from soil level

Burn Depth(m) = (Post fire height - Pre fire height)(m)



Figure 3. A visual representation of the box model method for calculating remaining perturbation concentration

Impulse Response Function- Derived from model simulations, expressed as fraction of initial concentration remaining, use for  $CO_2$ 

<b>Figure 2.</b> Example measurement of burn depth of <i>Dicranum</i>
Equation 2. Amount of biomass or carbon lost from combustion
based on soil characteristics of nearby unburned soil
Biomass or Carbon $\left(\frac{kg}{m^2}\right) = \frac{Dry Mass(kg)}{Volume(m^3)} \times$
Biomass or Carbon $(kg)$ , $\mathbf{D}$ , $\mathbf{D}$

 $\xrightarrow{s}$  × Burn Depth(m) Dry Mass (kg) **Equation 3.** Amount of biomass burned in the fire

Burned biomass (kg) = Biomass  $\left(\frac{kg}{m^2}\right) \times$  Fire area (m<sup>2</sup>)

Fire area = 726 km<sup>2</sup>, determined from Landstat dNBR, pixels were RCP4.5- moderate emissions reduction assigned severity based on magnitude of dNBR value

### Calculate greenhouse gases released and their radiative forcing

Model Inputs: year of burn, time after fire analyzed, scenario, dry biomass combusted Steps:

- Convert dry biomass to mass of gas species emitted using emissions factors
- Calculate the amount of that species in the atmosphere per year after fire
- Calculate radiative forcing from concentration of species in the atmosphere per year

Interpolation between GWPs, use for short-lived climate forcers

• Temporally explicit equations, use for  $CH_4$ ,  $O_3$ ,  $CO_2$  and  $N_2O$ • Atmospheric model outputs, use for aerosols

**Equation 4.** The radiative forcing for a perturbation concentration of  $CO_2$ .<sup>1</sup>

RF<sub>CO</sub>

## **Carbon loss and radiative forcings of gaseous emissions from tundra** wildfires during the Yukon Kuskokwim River Delta 2015 fire season

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### **Methods for calculating remaining perturbation** concentration:

Box Model-Assume one inflow and one outflow, use for  $CH_4$ ,  $O_3$ ,





Figure 4. The impulse response function derived from multiple model simulations for  $CO_2^1$ 

RCP scenarios (future emissions projections): Radiative forcing and atmospheric lifetime depend on ambient concentration of species.

- Historic- ambient concentration set at 2018 value
- RCP8.5- business as usual

## **Methods for calculating radiative forcing:**

$$N_{\rm CO_2}(\Delta N_{\rm CO_2}) = 5.35 \,\mathrm{W \,m^{-2} \,ln} \left( \frac{N_{\rm CO_2,0} + \Delta N_{\rm CO_2}}{N_{\rm CO_2,0}} \right)$$



Species Figure 5. The average burn depth per transect calculated for each vegetative marker (Eq. 1). Means between Dicranum and *Sphagnum* were comparable, but means between moss genera and tussock differed. All means across burn severity were significant for moss genera (p < 0.05) but not for tussocks. Radiative forcing remains positive after fires for long time horizons Figure 8. Cumulative mean radiative forcing at 20- and 80years post fire of main greenhouse gas species, short lived climate forcers, and aerosol effects for the RCP 8.5 scenario based on an average combustion value of 2.83  $\left(\frac{kg}{m^2}\right)$  of dry biomass

## Conclusions

# Acknowledgements

development.

## References

1. Joos, F., et al. (2013). Carbon dioxide and climate impulse response functions for the computation of greenhouse gas metrics: a multi-model analysis, Atmospheric Chemistry and Physics, 13, 2793-2825. 2. Ward, D. S, et al. (2012). The changing radiative forcing of fires: global model estimates for past, present and future, Atmospheric Chemistry and Physics, 12(22), 10857-10886.



# **Results -** The 2015 fire season released 1.32 Tg of carbon to the

Species

Burn depth and carbon loss varied across burn severity





Total w/o short lived climate forcers

Total w/ short lived climate forcers

NMVOC

N20

CO-

CH4-

–1 0 1 2 3 4 RF (w/m^2)

Indirect aerosol effects-

Direct aerosol effects





• Estimating burn depth was consistent across moss genera, but varied between mosses and tussock. Difference in burn depth across moss genera and tussocks and tussock burn depth across transect length is suggestive of altered microclimate by tussock grasses that affect flammability. Bulk density measurements also do not agree between moss genera and tussocks. Further research is needed to identify plant proxies for estimating plant and soil loss with fire.

Radiative forcing persists over long timescales and varies by gas and through time. Tundra fires have a significant warming effect for which should be accounted with accurate combustion measurements and radiative forcing models.

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Figure 6. The average carbon

released per transect calculated for each vegetative marker using Eq. 3. Inter-species means agreed, but means across severity varied (p <0.01).

### Radiative forcing varied by gas and

Carbon Dioxide Carbon Monoxide Direct Indirect

Figure 9. Cumulative radiative forcing of greenhouse gas species 100 years post fire for the RCP 8.5 scenario based on an average combustion value of 2.83  $\left(\frac{kg}{m^2}\right)$  of dry biomass